

# Tsunami wave energy

Denys Dutykh<sup>1</sup> Frédéric Dias<sup>1</sup>

<sup>1</sup>Ecole Normale Supérieure de Cachan,  
Centre de Mathématiques et de Leurs Applications,  
LRC Méso CMLA/CEA DAM IdF

## Nonlinear Coherent Wave Structures in Fluids



# Outline

- 1 Physical context
  - Motivation
  - Tsunami generation analysis
- 2 Energy equation derivation
- 3 Numerical results
- 4 Conclusions

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# Energy of tsunamis

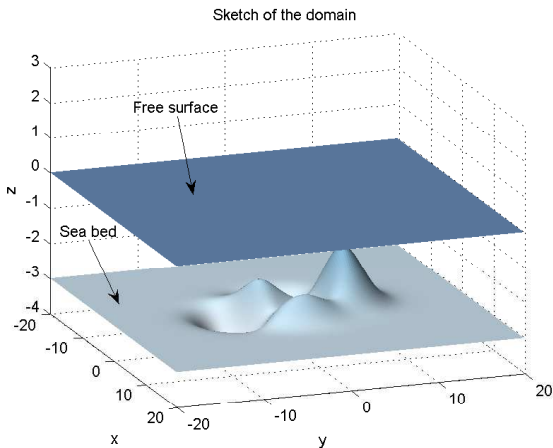
Existing literature (this list is not exhaustive)

- 1 S. Tinti & E. Bortolucci. *Energy of Water Waves Induced by Submarine Landslides*. Pure appl. geophys. 157 (2000)
- 2 Z. Kowalik et al. *The Tsunami of 26 December 2004 : Numerical Modeling and Energy Considerations*. 22nd IUGG International Tsunami Symposium (2005)
- 3 T. Murty et al. *Leakage of the Indian Ocean Tsunami Energy into the Atlantic and Pacific Ocean*. J. Canadian Association of Exploration Geophysicists (2005)
- 4 A. Velichko et al. *Amplitude-energy characteristics of tsunami waves for various types of seismic sources generating them*. Physical Oceanography (2002)

# Traditional approach

## Approaches to generation

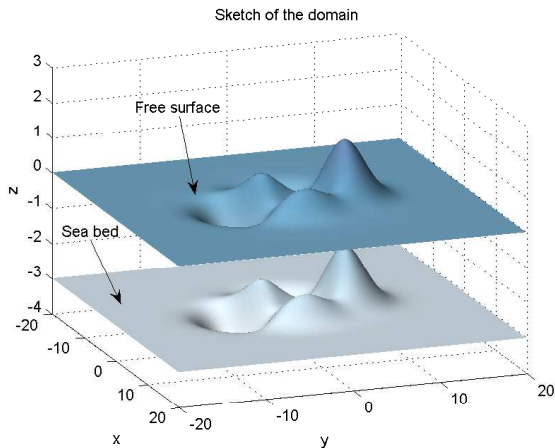
Put coseismic displacements directly on the free surface and let it propagate :



# Traditional approach

## Approaches to generation

Put coseismic displacements directly on the free surface and let it propagate :



# Tsunami generation process

Different scenarios of bottom motion

We consider the following situations :

- 1 Passive generation
  - $\Rightarrow$  Initial Value Problem (IVP)
- 2 Instantaneous seabed deformation

$$h(\vec{x}, t) = h_0(\vec{x}) + \mathcal{H}(t)\zeta(\vec{x})$$

- 3 Exponential scenario

$$h(\vec{x}, t) = h_0(\vec{x}) + (1 - e^{-\alpha t})\zeta(\vec{x})$$

## Remark

Passive generation is not equivalent to instantaneous seabed motion

# Analytical solutions

## Linearized water waves : Cauchy-Poisson problem

- Free surface elevation

$$\eta(\vec{x}, t) = \frac{1}{(2\pi)^2} \iint_{\mathbb{R}^2} \frac{\widehat{\zeta}(\vec{k}) e^{i\vec{k} \cdot \vec{x}}}{\cosh(|\vec{k}|h)} \frac{1}{2\pi i} \int_{\mu-i\infty}^{\mu+i\infty} \frac{s^2 \mathbf{T}(s)}{s^2 + \omega^2} e^{st} ds d\vec{k}$$

- Velocity potential

$$\hat{\phi}(\vec{x}, z, t) = \frac{gs\widehat{\zeta}(\vec{k})\mathbf{T}(s)}{\cosh(|\vec{k}|h)(s^2 + \omega^2)} \left( \frac{s^2}{g|\vec{k}|} \sinh(|\vec{k}|z) - \cosh(|\vec{k}|z) \right)$$

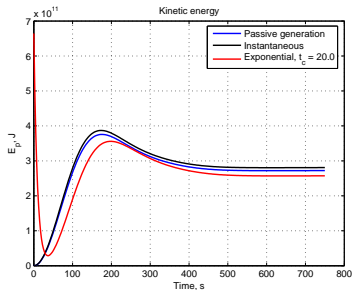
- Compute the energies

$$K = \frac{1}{2} \int_{-h}^{\eta} \iint_{\mathbb{R}^2} |\nabla \phi|^2 d\vec{x} dz \quad \Pi = \frac{g}{2} \iint_{\mathbb{R}^2} \eta^2 d\vec{x}$$

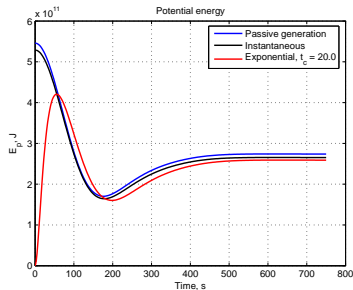


# Energy analysis of generation process - I

## Kinetic and potential energies



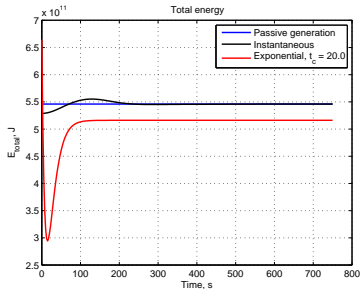
(a) Kinetic energy



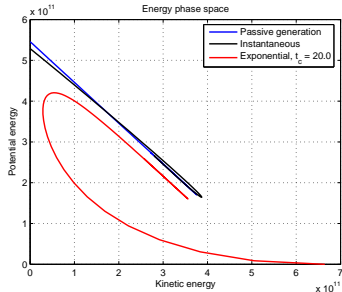
(b) Potential energy

# Energy analysis of generation process - II

## Total energy and phase space



(c) Total energy



(d) Energy phase space

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# Energy equation derivation for long waves

Starting point : Euler equations with free surface

After integrating energy equation over the depth :

$$E_t + \nabla \cdot \Phi + P = 0$$

- Total energy

$$E = \int_{-h}^{\eta} \left( \frac{1}{2} \rho (|\vec{u}|^2 + w^2) + \rho g z \right) dz$$

- Energy flux

$$\Phi = \int_{-h}^{\eta} \vec{u} \left( \frac{1}{2} \rho (|\vec{u}|^2 + w^2) + p + \rho g z \right) dz$$

- Energy input term

$$P = p_s \eta_t + p_b h_t$$

# Energy equation derivation for long waves

Switching to dimensionless quantities

After integrating energy equation over the depth :

$$E_t + \varepsilon \nabla \cdot \Phi + \varepsilon P = 0$$

- Total energy

$$E = \int_{-h}^{\varepsilon \eta} \left( \frac{\varepsilon}{2} (\varepsilon |\vec{u}|^2 + \frac{\varepsilon}{\mu^2} w^2) + z \right) dz$$

- Energy flux

$$\Phi = \int_{-h}^{\varepsilon \eta} \vec{u} \left( \frac{\varepsilon}{2} (\varepsilon |\vec{u}|^2 + \frac{\varepsilon}{\mu^2} w^2) + \varepsilon p + z \right) dz$$

- Energy input term

$$P = \varepsilon p_b \zeta_t$$

# Governing equations - I

## NSW2E : Nonlinear Shallow Water Equations with Energy

- Mass conservation

$$H_t + \nabla \cdot (H\vec{u}) = 0$$

- Momentum conservation

$$(H\vec{u})_t + \nabla \cdot (H\vec{u} \otimes \vec{u} + \frac{g}{2}H^2) = gH\nabla h$$

- Energy equation

$$(HE)_t + \nabla \cdot ((HE + \rho\frac{g}{2}H^2)\vec{u}) + \rho gH\zeta_t = 0$$

# Governing equations - II

## Boussinesq equations with energy

- Mass conservation

$$H_t + \nabla \cdot (H\vec{u}) = 0$$

- Momentum conservation

$$(H\vec{u})_t + \nabla \cdot \left( H\vec{u} \otimes \vec{u} + \frac{g}{2}H^2 \right) + \left( \frac{h^3}{6} \nabla (\nabla \cdot \left( \frac{H\vec{u}}{d} \right)) - \frac{h^2}{2} \nabla (\nabla \cdot (H\vec{u})) \right)_t = gH\nabla h$$

- Energy equation

$$(HE)_t + \nabla \cdot \left( (HE + \rho \frac{g}{2}H^2)\vec{u} \right) + \mathcal{O}(\varepsilon\mu^2) + \rho gH\zeta_t = 0$$

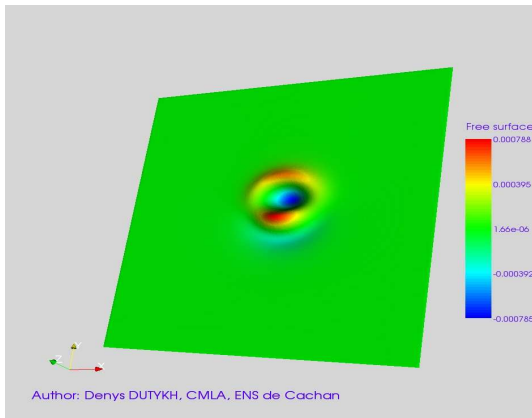
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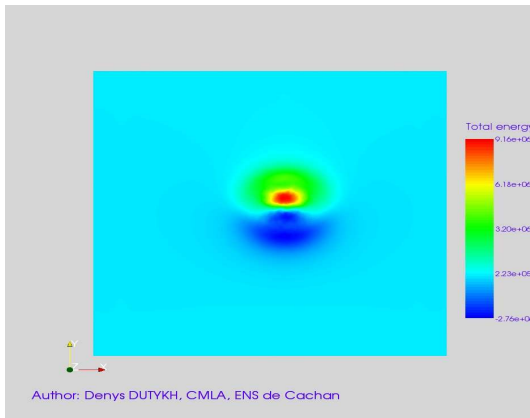
# Active generation of a tsunami wave - I

Free surface elevation



# Active generation of a tsunami wave - II

Total energy density plot



# Energy reconstruction

Computation of the total energy from NSWE solutions

- 1 We solve NSWE  $\Rightarrow (\eta, h, \vec{u})$
- 2 Using this information, we approximate the energy density :

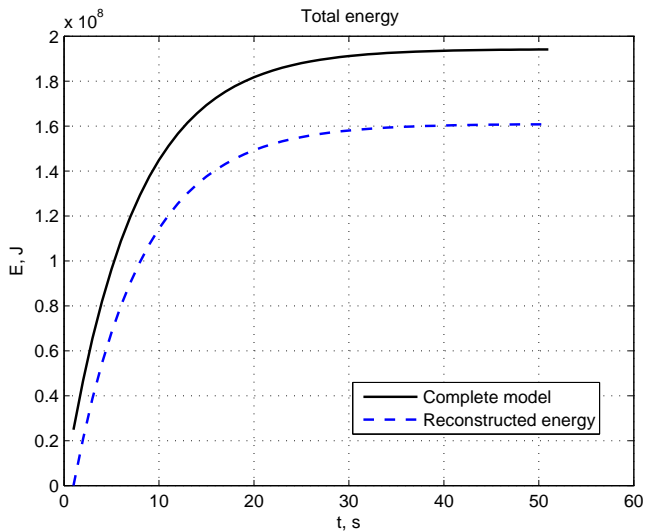
$$E \approx (h + \eta) \left( \frac{1}{2} \rho |\vec{u}|^2 + \frac{1}{2} \rho g (\eta^2 - h^2) \right)$$

- 3 To obtain the total energy, we integrate over the domain  $\Omega$  :

$$\mathbb{E}(t) = \iint_{\Omega} E(\vec{x}, t) d\vec{x}$$

# Total energy

Comparison between different ways of energy computation



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# Main conclusions

## Preliminary results

- Different approaches for tsunami generation are not equivalent from energy viewpoint
  - : Passive generation overestimates tsunami energy
  - + : All lead to the equipartition between  $K$  and  $\Pi$
- Energy propagation is **nondispersive**
- For long waves solving the additional energy equation is not equivalent to energy reconstruction a posteriori

Thank you for your attention !

<http://www.cmla.ens-cachan.fr/~dutykh>